

# What controls the biological pump in the ocean ?

## Implications for future warming and ocean acidification

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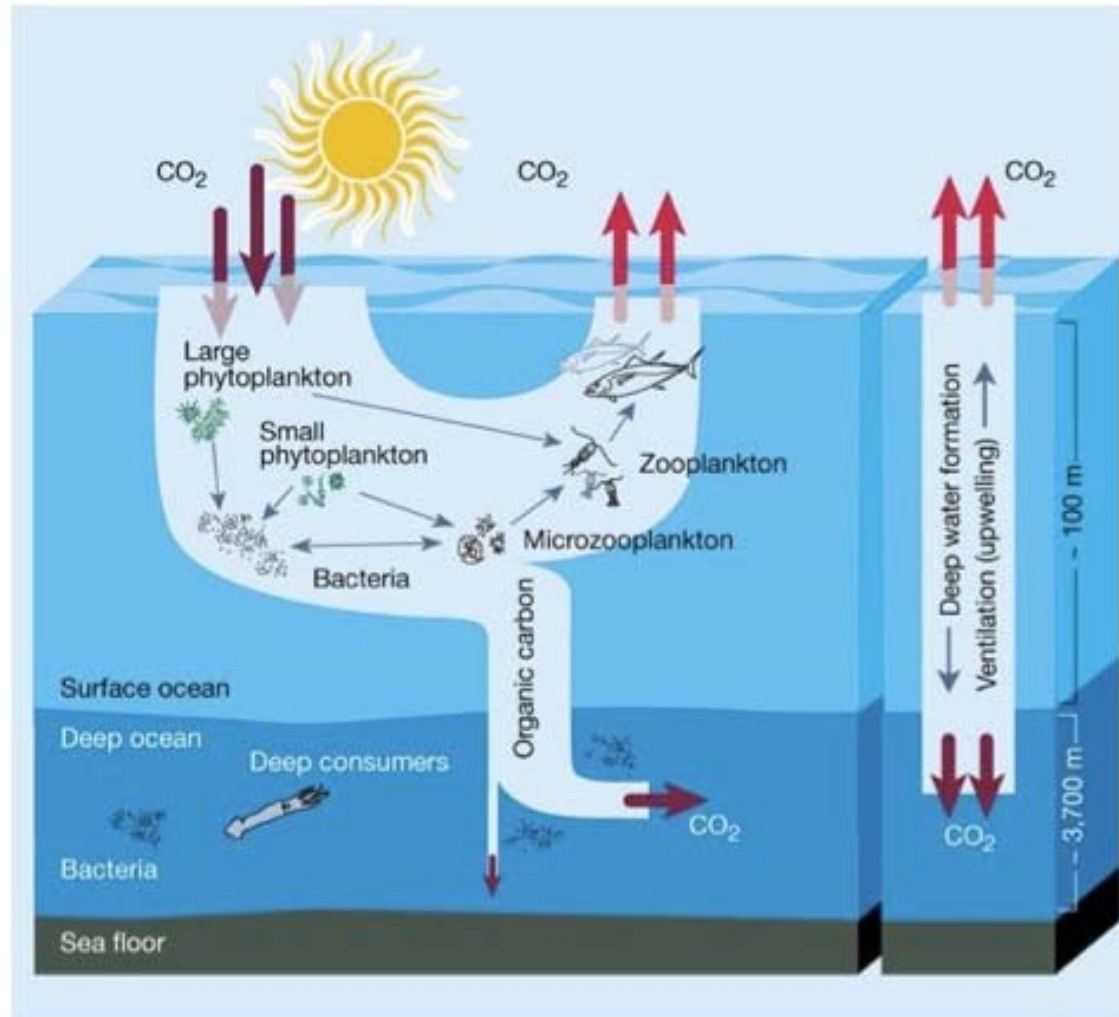
<sup>4</sup>University of Leeds, UK



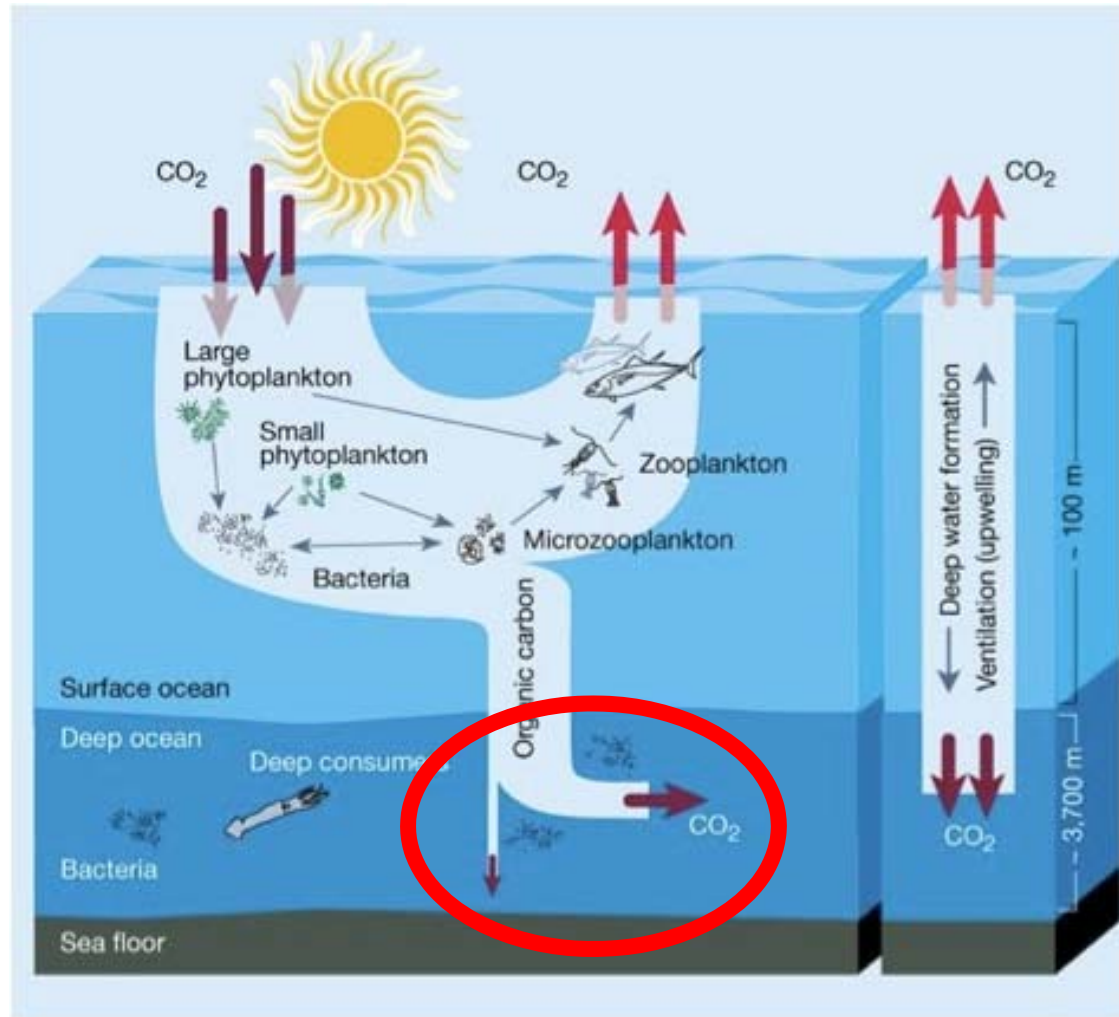
UK Ocean Acidification  
Research Programme



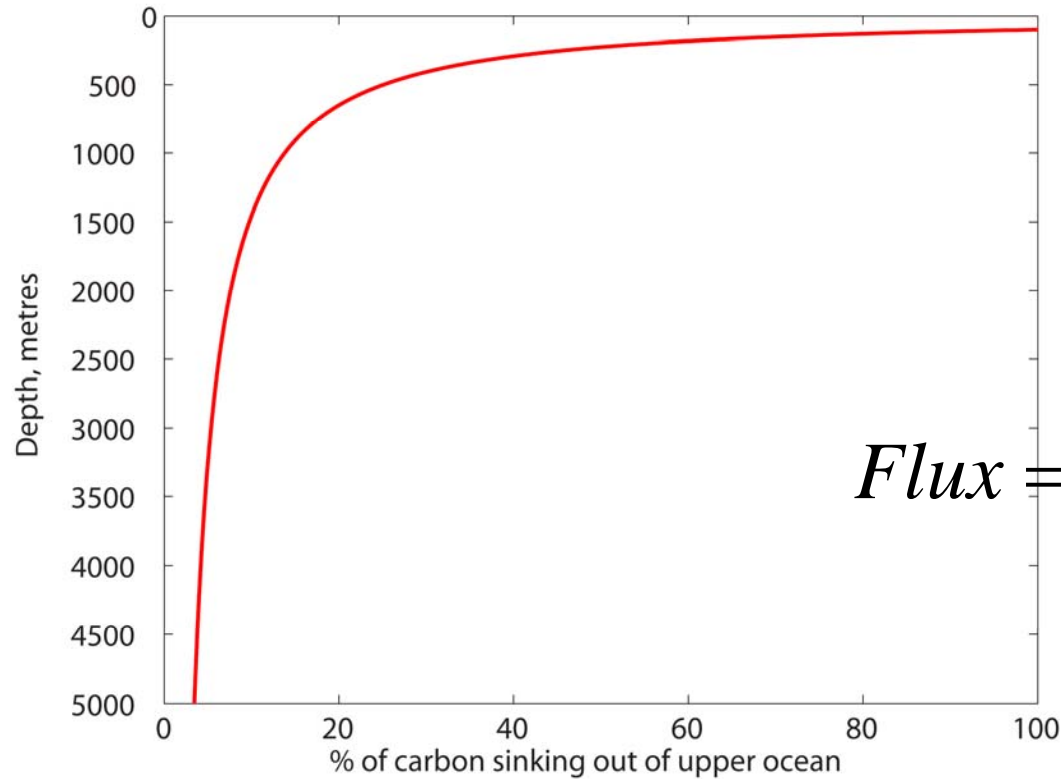
# 🌞 The biological pump



# 🌞 The biological pump



# 🌿 How is Organic Carbon transferred to depth?

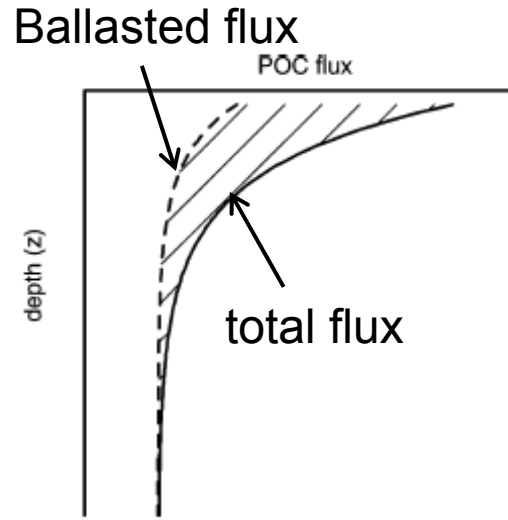


Martin et al. (1987)

$$Flux = Export \cdot \left( \frac{z}{100} \right)^{-b}$$

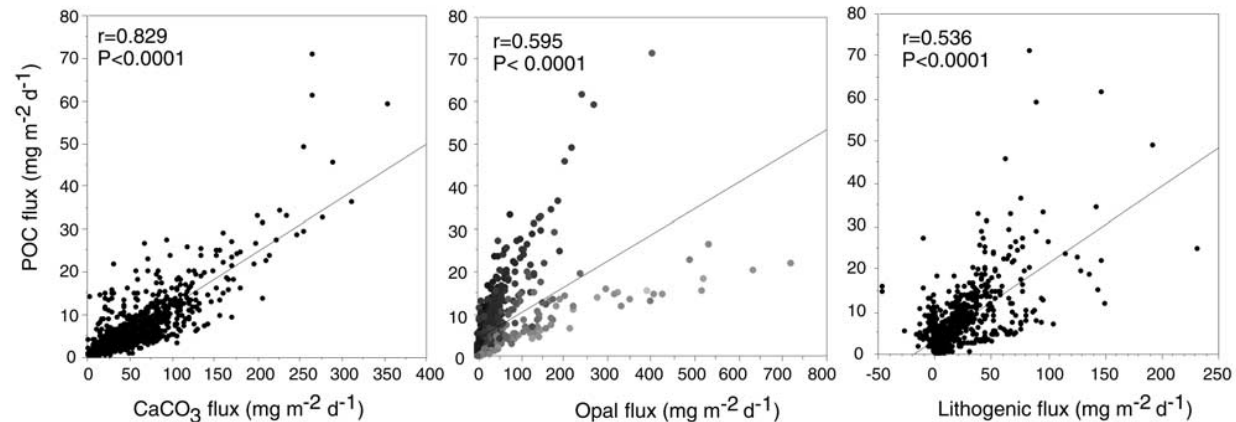
- Martin's flux curve and original  $b$  value based on 9 flux profiles in North Pacific;  $b = 0.898$
- $b$  initially assumed to be constant

# Ballast hypothesis



Armstrong et al. 2002

- Biominerals (particularly  $\text{CaCO}_3$ ), increase density of sinking particles and/or provide protection against degradation



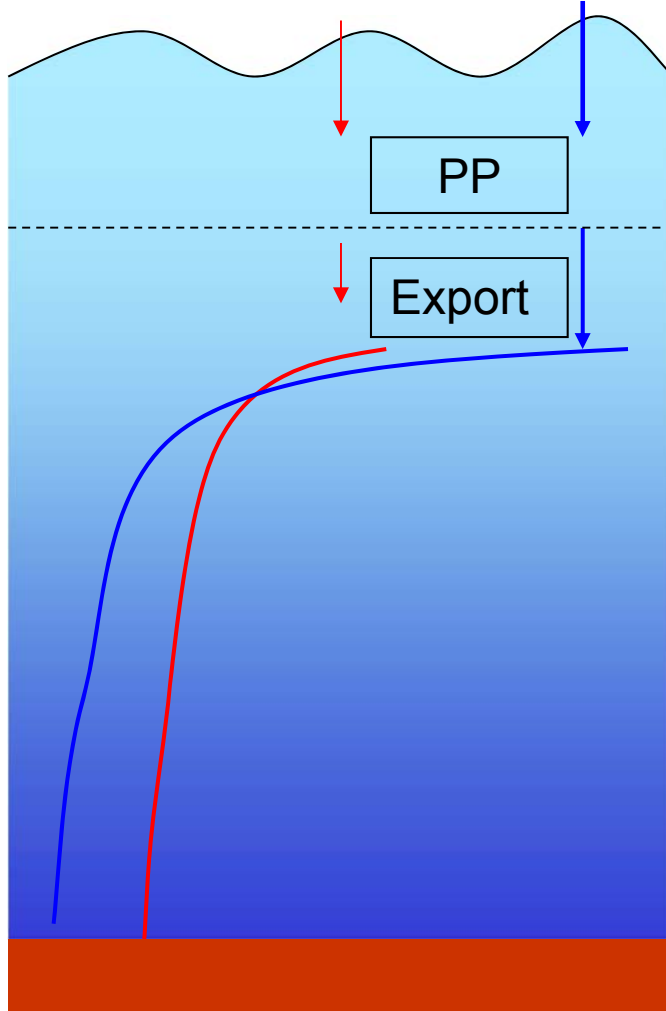
$$F_{POC} = \beta_1 \cdot F_{\text{CaCO}_3} + \beta_2 \cdot F_{\text{opal}} + \beta_3 \cdot F_{\text{litho}}$$

Klaas and Archer  
(2002)

POC: Particulate Organic Carbon

# 🌿 Ecosystem function

More C delivered to mesopelagic at high latitudes →  
food web structure?



High export efficiency

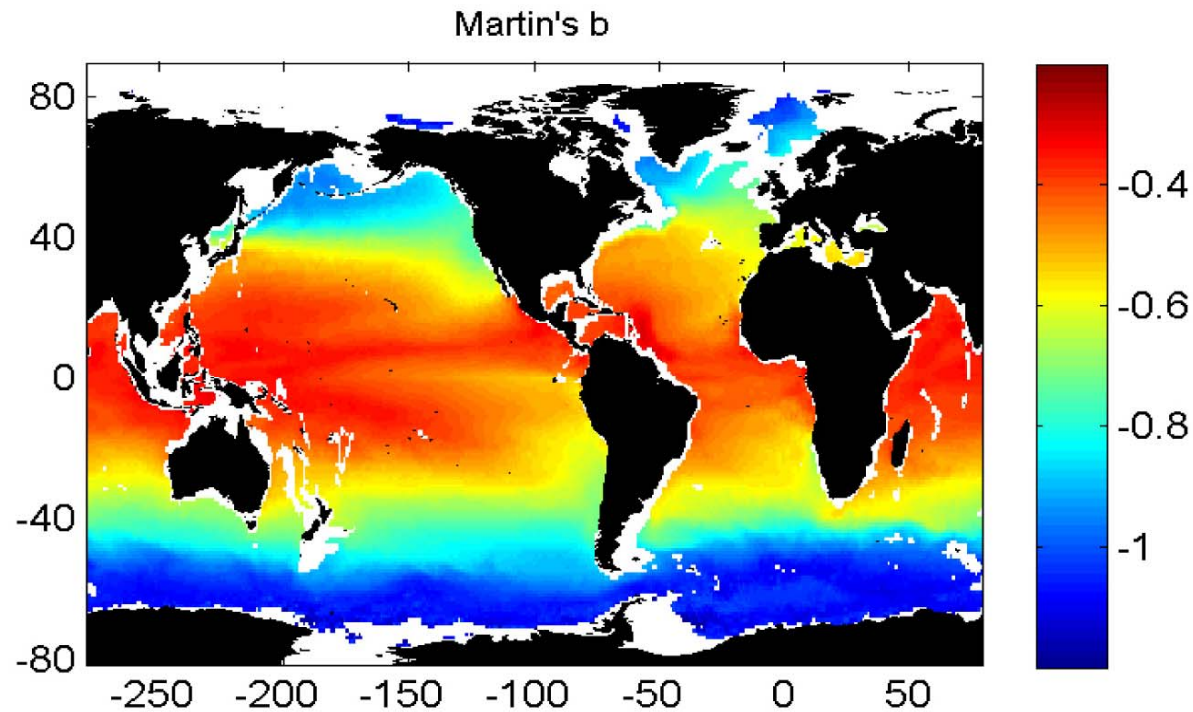
Low export efficiency

Low transfer efficiency

High transfer efficiency

🌟 Temperature dependent transfer efficiency

$$Flux = Export \cdot \left( \frac{z}{100} \right)^{-b} \quad b = (0.024 * SST) - 1.06$$



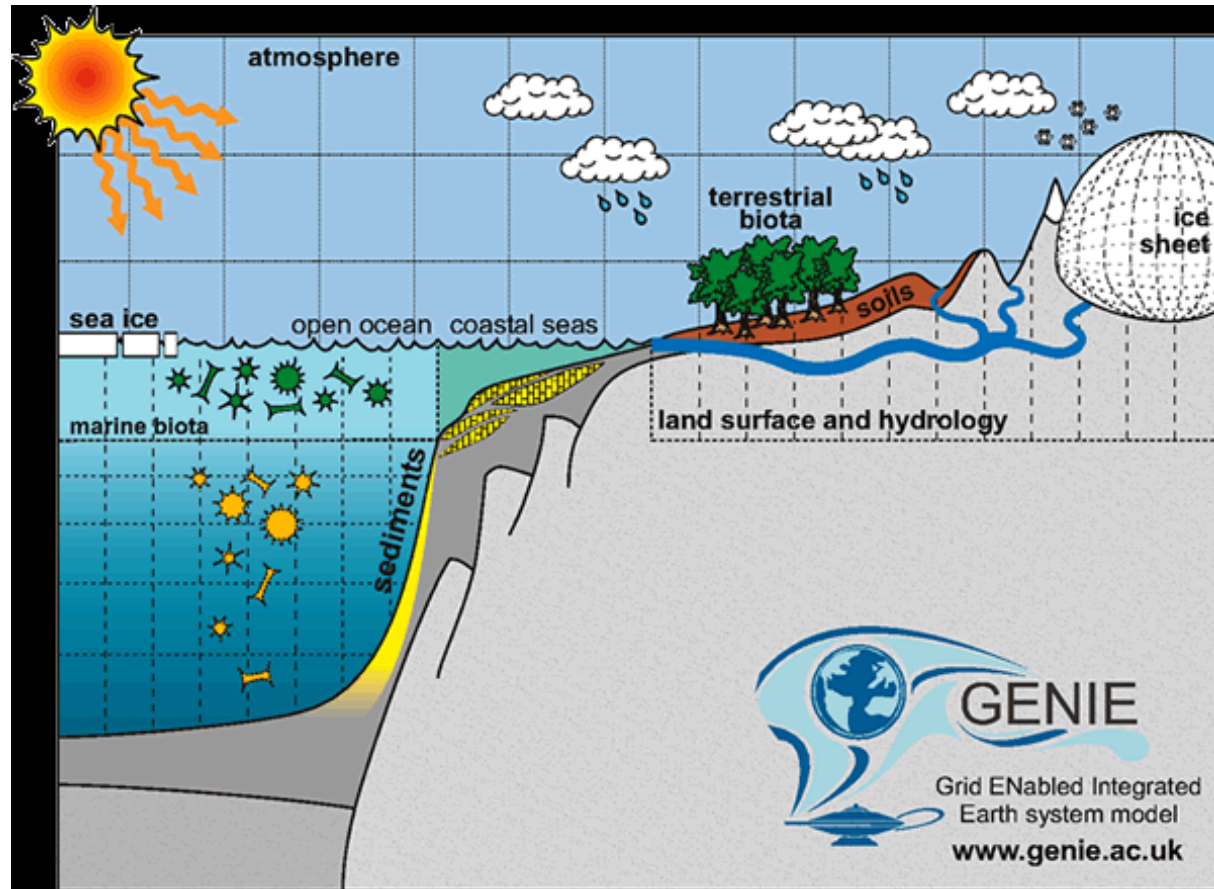
Henson et al., 2012

# 🌞 Earth System Model: cGenie

Represents biogenic biogeochemical fluxes

Organic Matter Export is a function of :

- Nutrient ( $\text{PO}_4$ , Fe),
  - light
  - temperature
  - mixed layer depth
- (Death et al., in prep.)  
based on Doney et al., 2006





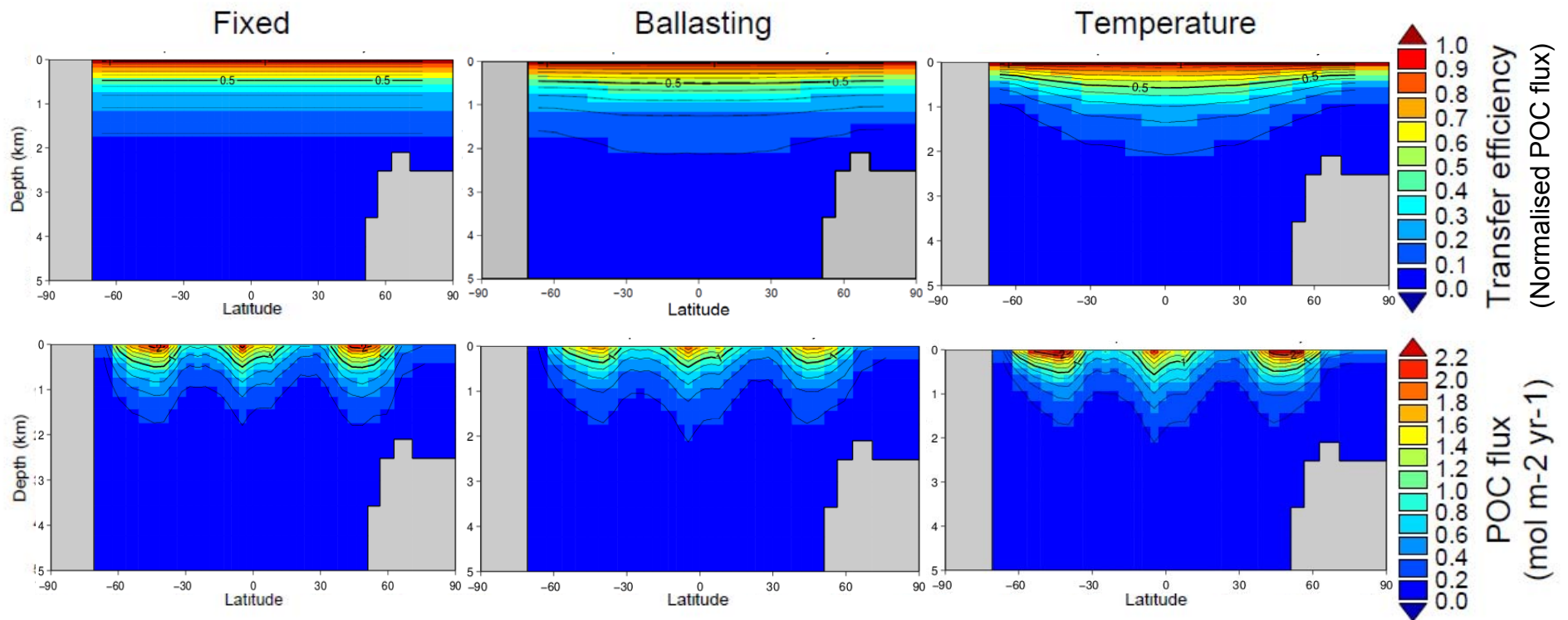
## ✦ Modelling transfer efficiency

- Standard: 2 fractions of organic matter with exponential decay  
Fixed everywhere and constant through time
- Ballasting: one fraction depends on  $\text{CaCO}_3$  flux
- Henson et al. (2012, GBC):  
Exponential decay dependent on sea surface temperatures

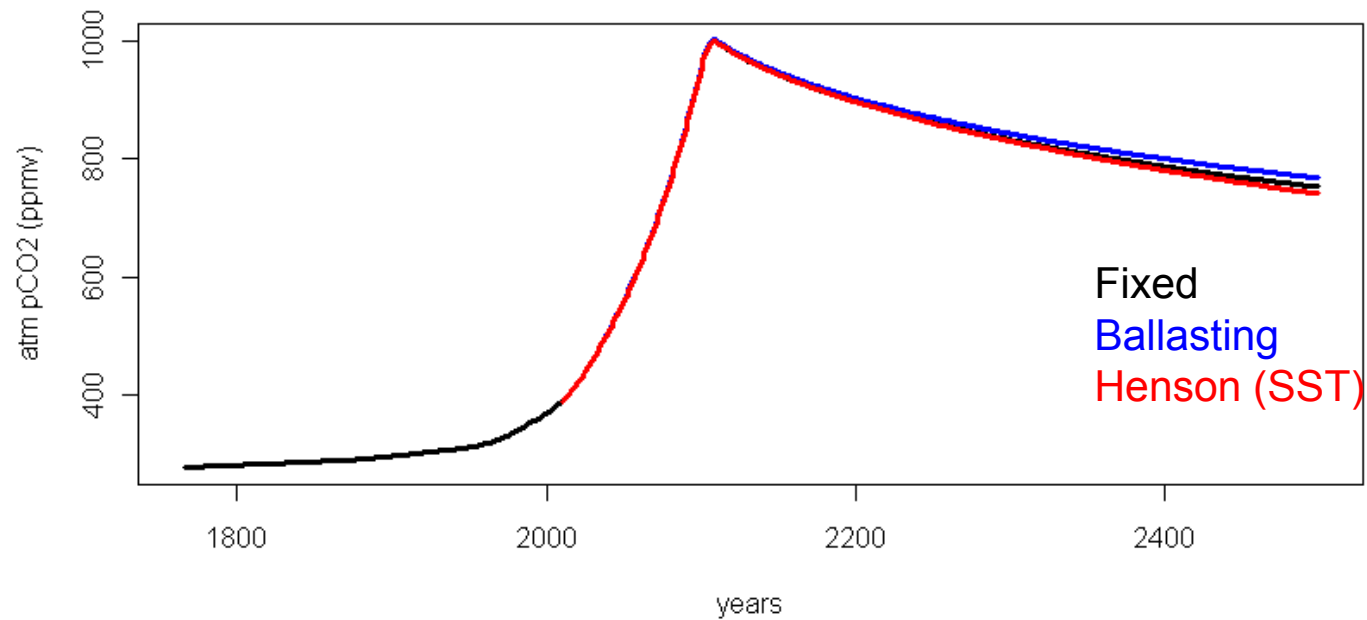
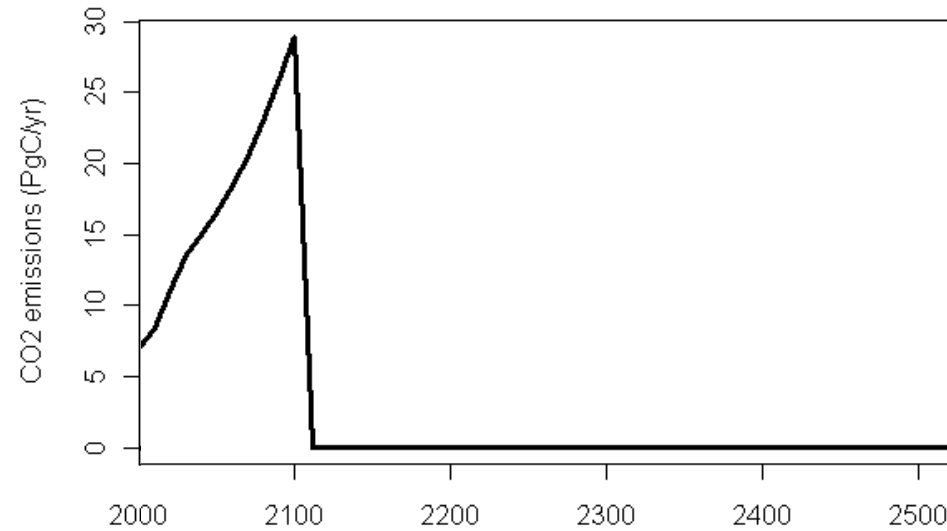
## Recalibrating cGenie

- Changing the POC flux profile redistributes nutrients in the ocean
  - Need to recalibrate model
- We optimised :
  - biological uptake timescale (control on POC export)
  - remineralisation profile parameters
- Selected simulations with smallest mean squared error for 3D annual  $\text{PO}_4$  and  $\text{O}_2$  (World ocean atlas 2009).

# 🔥 Profiles of transfer efficiency

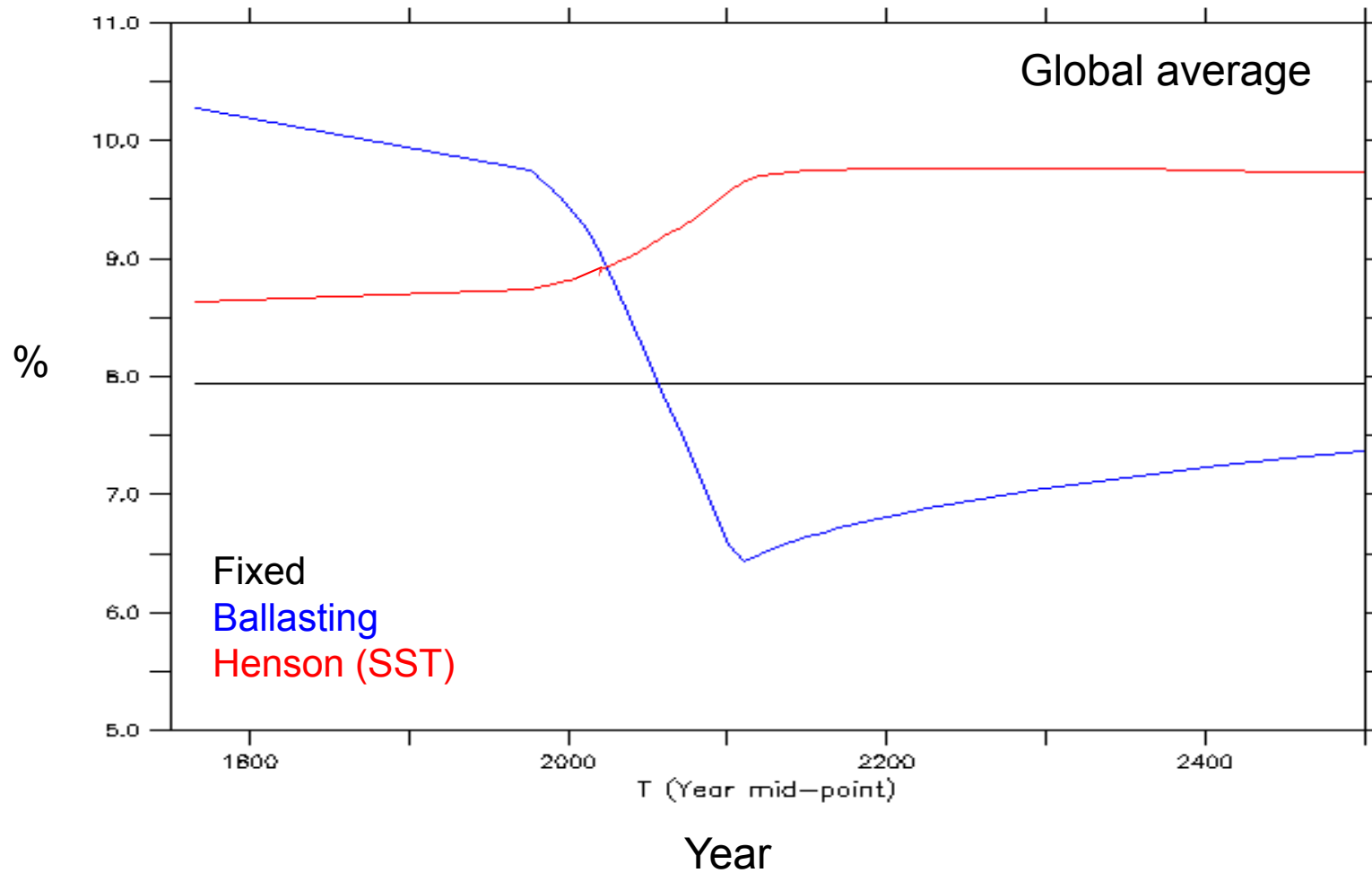


# 🌞 Future projections



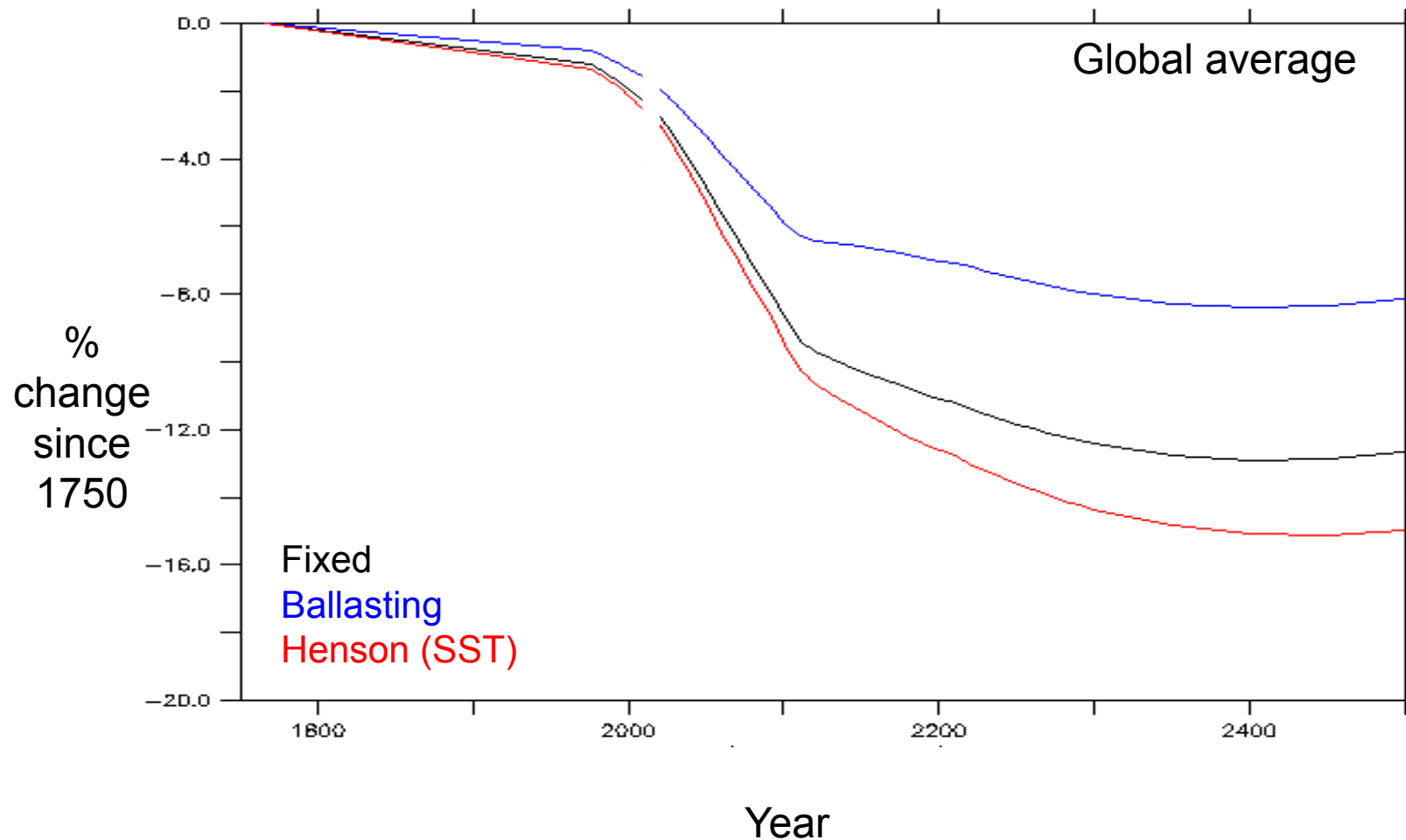
# Transfer efficiency

$$T_{\text{eff}} = 100 * F_{\text{POC}}(2000\text{m}) / \text{Export}_{\text{POC}} ; \text{Export}_{\text{POC}} = F_{\text{POC}}(100\text{m})$$



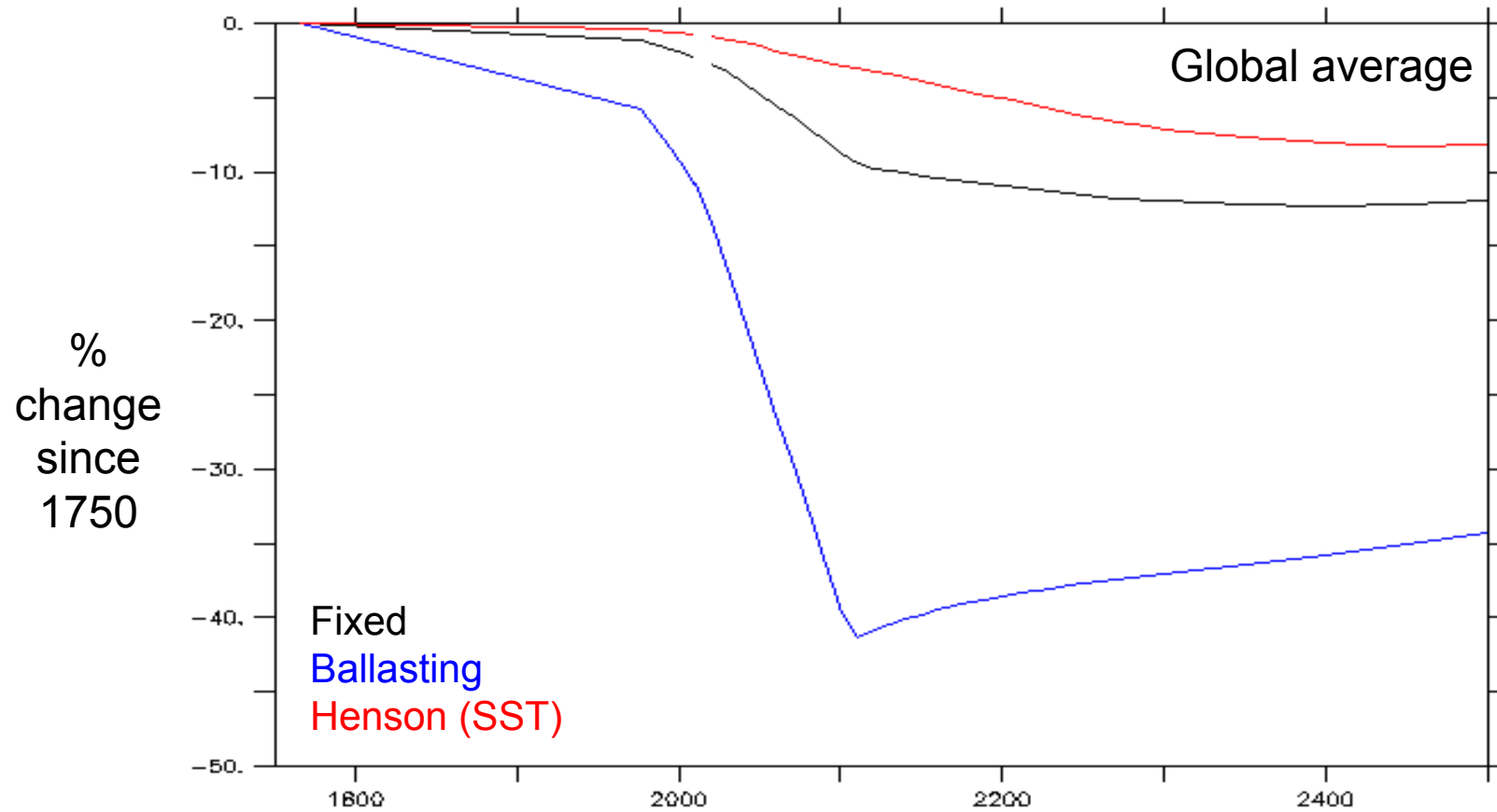
# 🌿 Feedback on export decrease

$$\text{Export}_{\text{POC}} = F_{\text{POC}}(100\text{m})$$



# 🌿 Change in deep POC flux

$F_{\text{POC}}(2000\text{m})$



# Conclusions

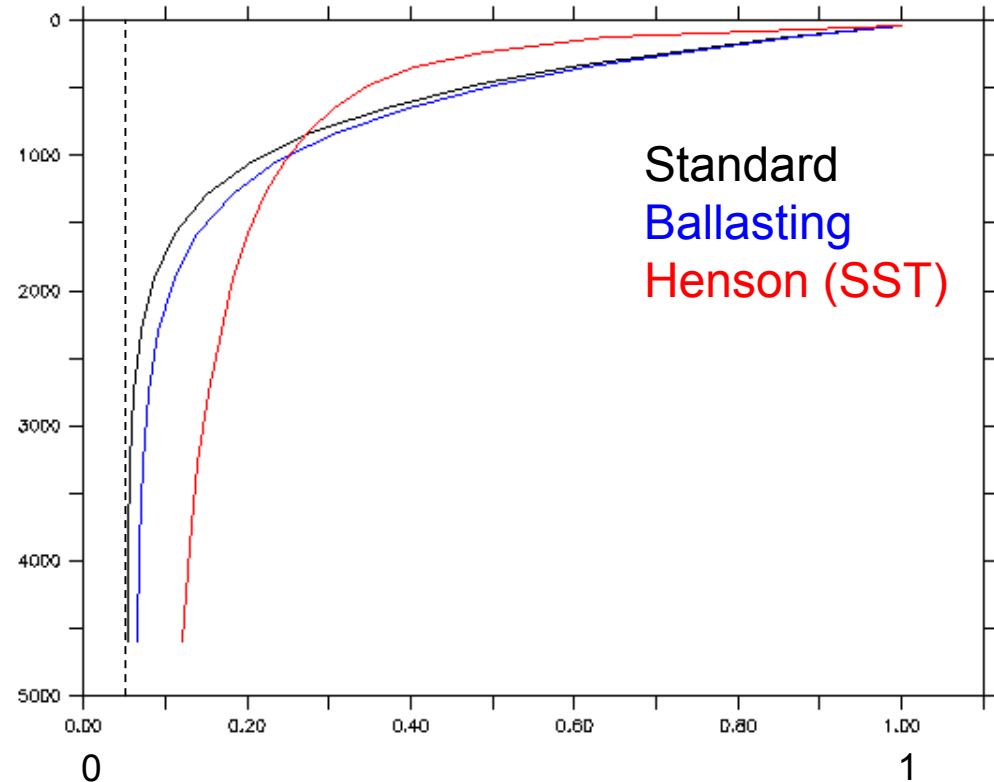
- Ballast hypothesis → decrease future transfer efficiency
- SST parameterisation → increased future transfer efficiency
- Different feedback on export decrease
- Effect on Nutrient distribution and O<sub>2</sub> min zones, but uncertainties in biological pump:
  - Amount of export
  - Control on organic matter degradation/recycling
  - Transfer of organic matter to deep ocean



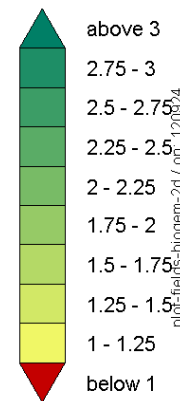
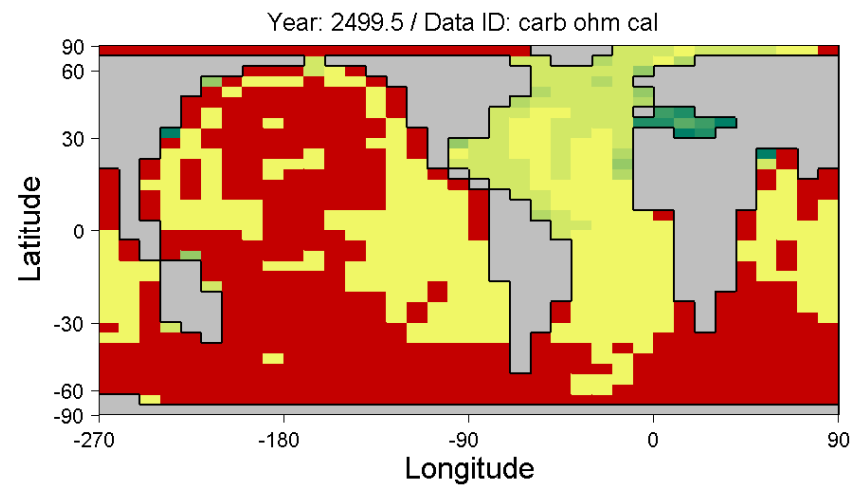
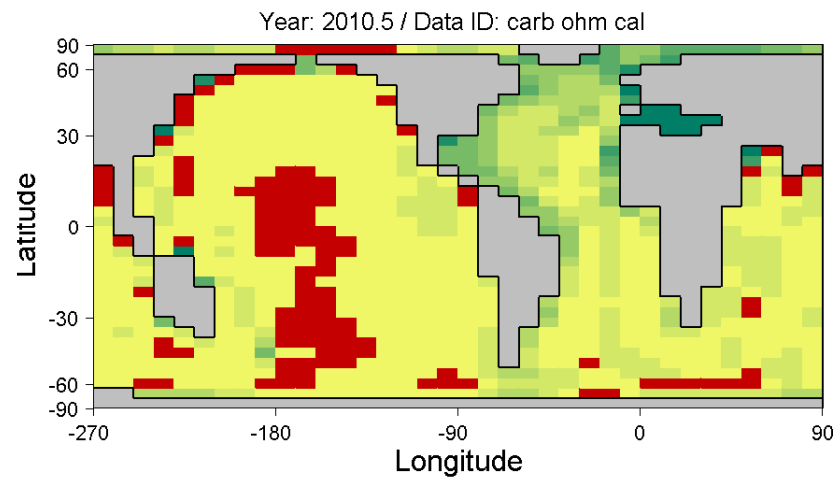


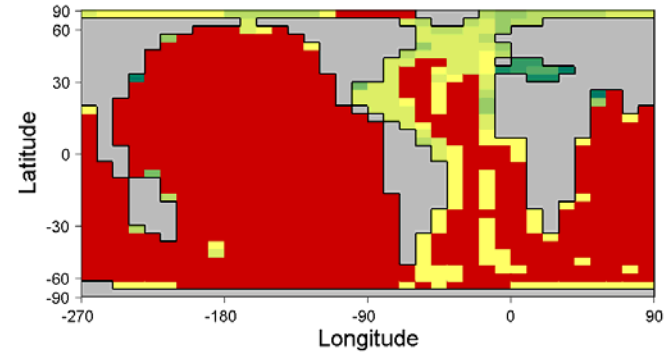
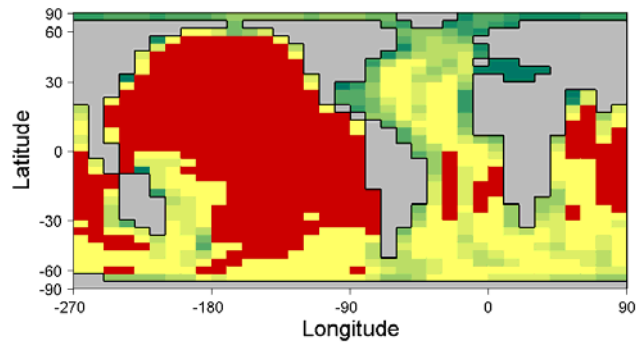
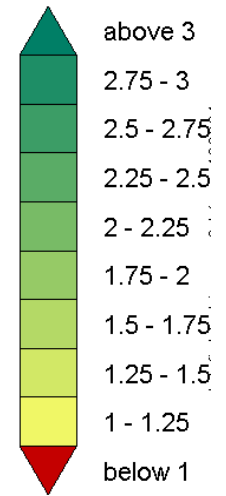
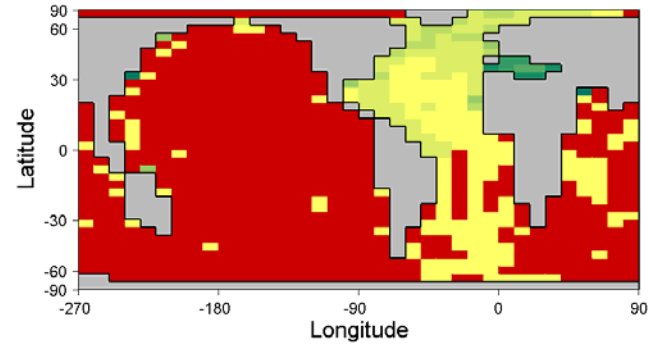
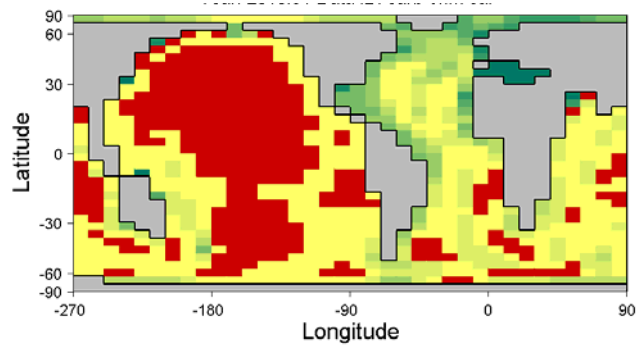
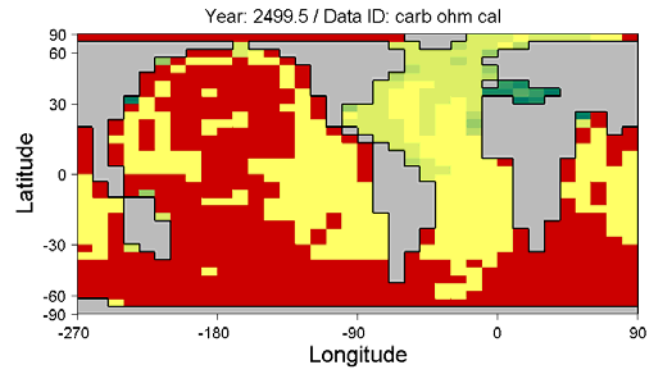
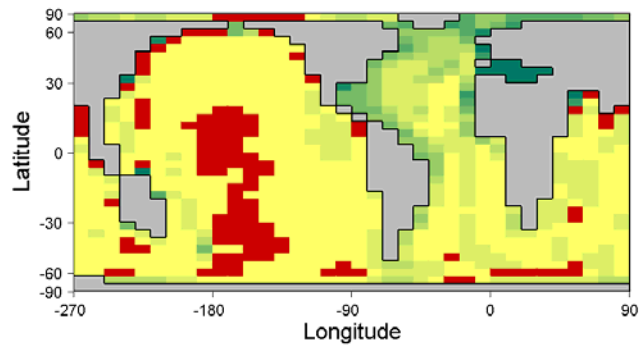
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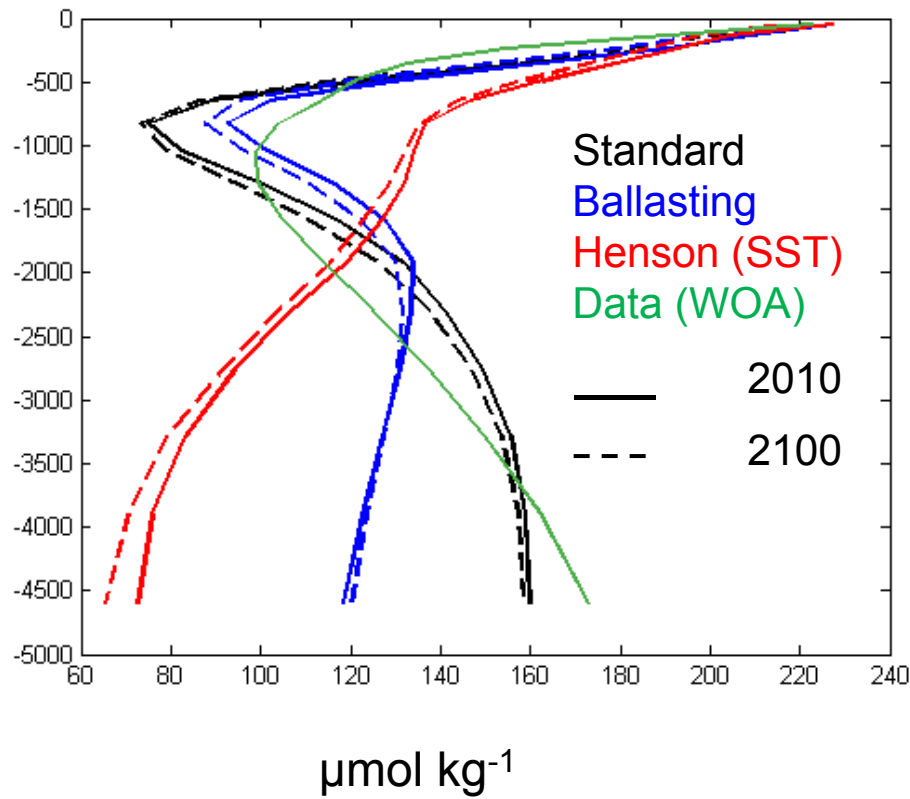
POC (particulate Organic Carbon) flux  
Normalised by Export (global average)



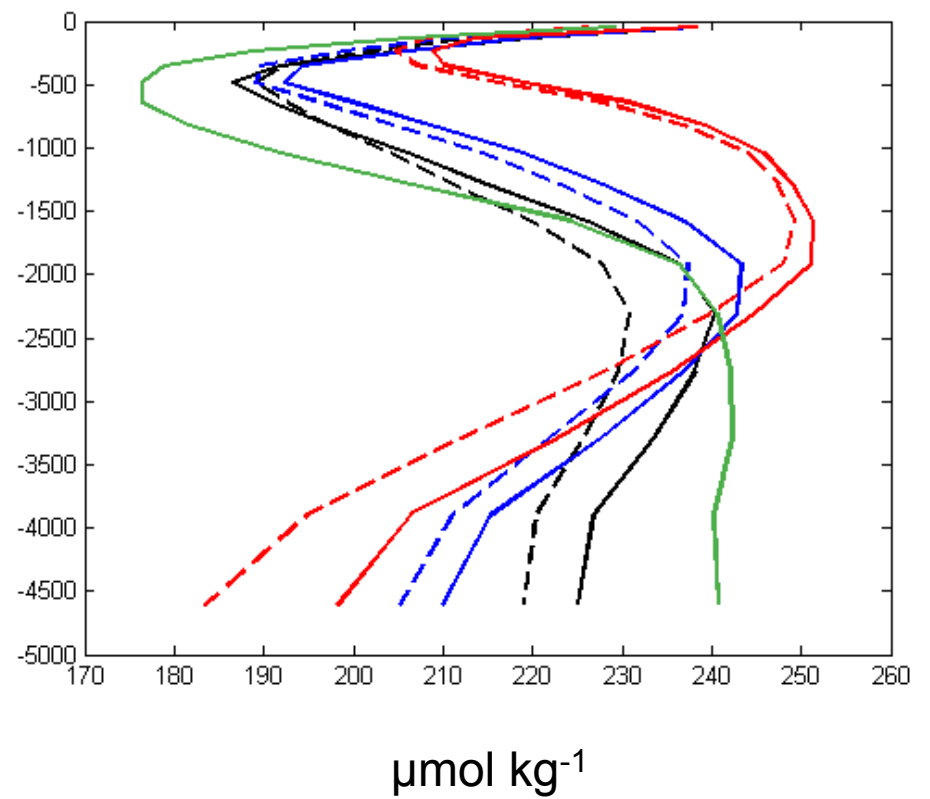


# 🌊 Dissolved Oxygen profiles

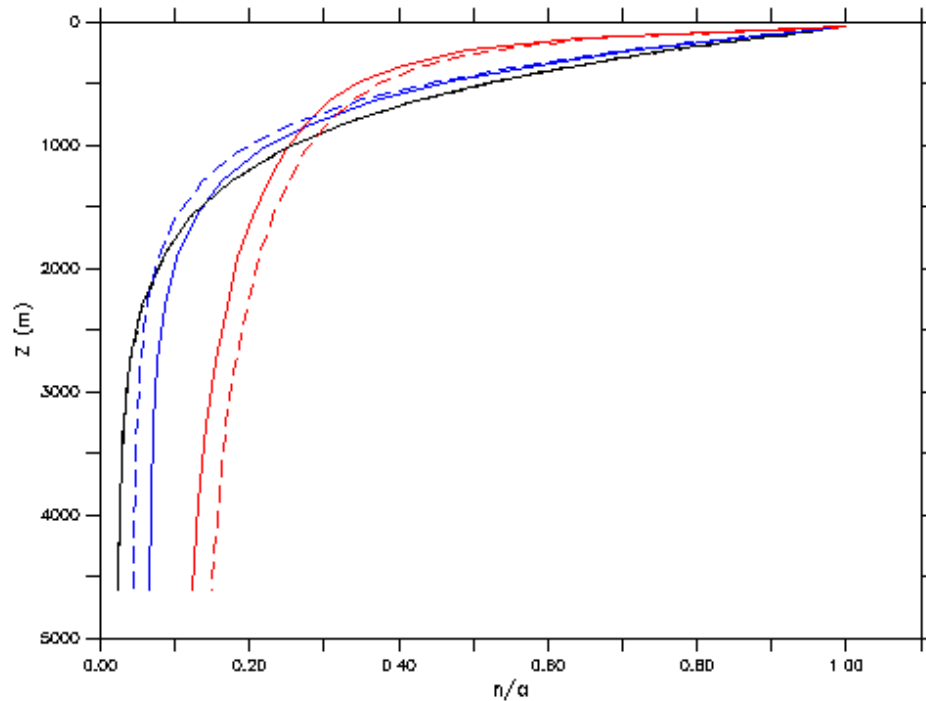
## Pacific



## Atlantic



# 🌋 Changes in POC profiles

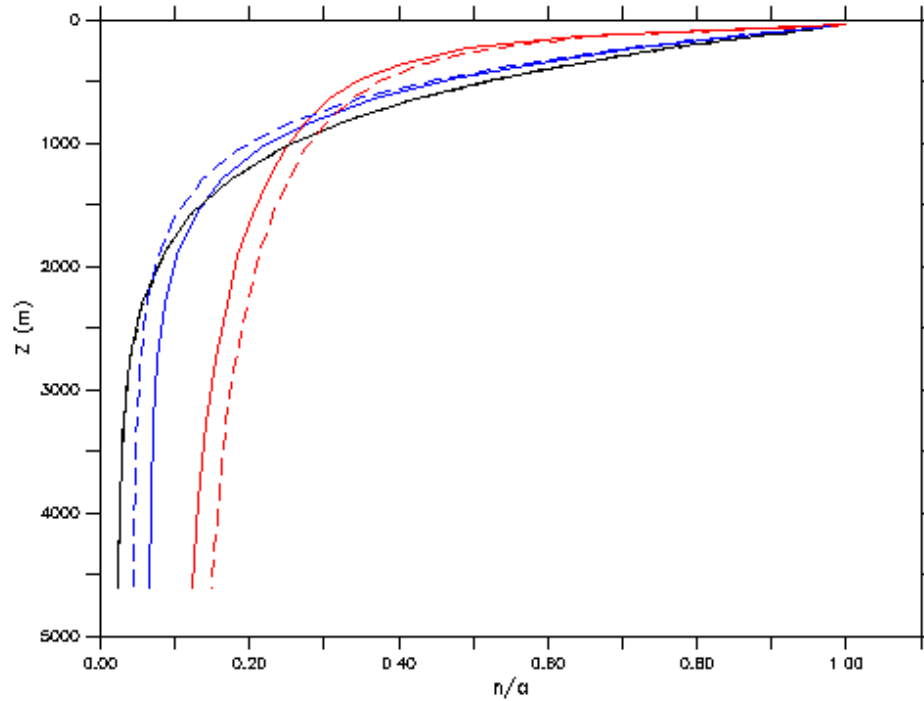


export-normalized particulate flux - POC (n/a)

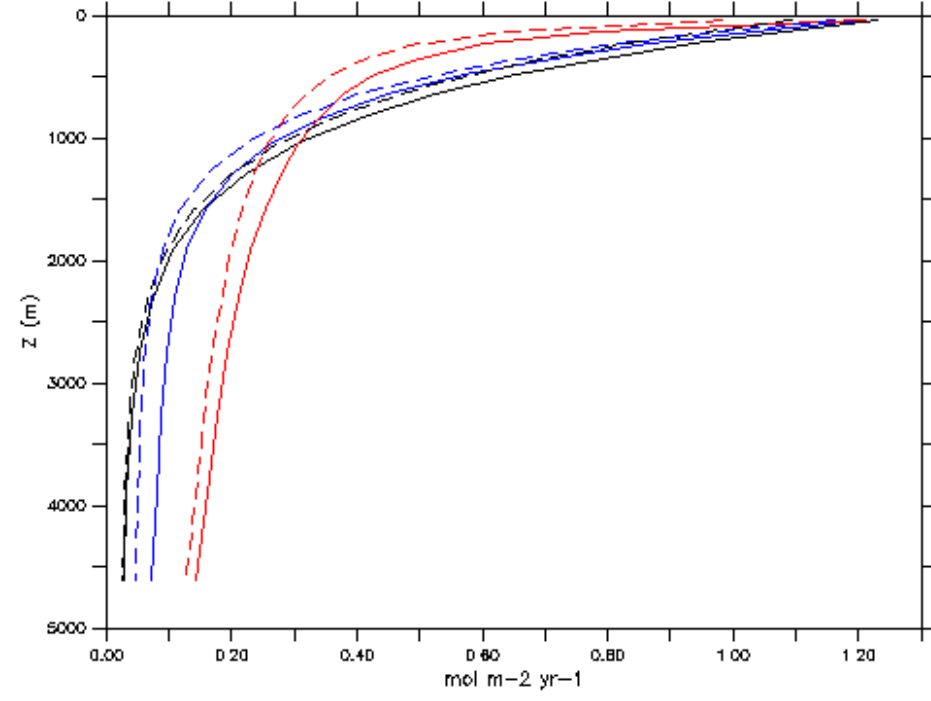
Standard	—	2010
Ballasting	—	
Henson (SST)	- - -	2500

# 🌊 Changes in POC profiles

Decreased export  
←



export-normalized particulate flux - POC (n/a)



particulate flux - POC (mol m<sup>-2</sup> yr<sup>-1</sup>)

Standard	—	2010
Ballasting	—	2010
Henson (SST)	- - -	2500

